

Evaluation of directionality in physics-based ground motion simulations of strike-slip earthquakes

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Recent advances in engineering seismology and computing power have led to the development of realistic physics-based ground motion simulations. The use of these simulations for engineering applications requires that the ground motions exhibit characteristics consistent with those of recorded ground motions. While several studies have evaluated the intensity, frequency content, duration, and other characteristics of ground motion simulations for their possible use in engineering applications, few have focused on directionality. This paper evaluates directionality in the response of single-degree-of-freedom oscillators when subjected to CyberShake Study 15.12 simulated ground motions from strike-slip earthquakes, by carefully comparing it to the directionality in recorded ground motions having the same style of faulting. Physics-based ground motion simulations at 334 stations from 5 different rupture variations in two large-magnitude earthquake scenarios on the Elsinore fault are evaluated. The orientation of maximum oscillator response and its spatial distribution are studied for each rupture. The orientation of maximum oscillator response is found to occur systematically close to the epicentral transverse orientation at all rupture distances, consistent with recent findings for ground motions recorded during strike-slip earthquakes. Additionally, the orientation-specific spectral accelerations, when computed as a function of angular distance from the epicentral transverse orientation, are found to exhibit a variation consistent with the overall trend observed in records from the NGA-West2 database. However, the level of polarization at short periods in the simulated hybrid broadband waveforms used in this study is larger than that in recorded ground motions.

INTRODUCTION

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27 Advances in seismology and computational power in the past few decades have led to the
28 development of sophisticated physics-based ground motion simulations (e.g., Graves and
29 Pitarka, 2010; Mai et al., 2010). Current design codes in the United States (ASCE,
30 American Society of Civil Engineers, 2022) permit the use of physics-based ground
31 motions as possible inputs for nonlinear response history analyses in the absence of suitable
32 recorded ground motions. Using simulated ground motions is especially beneficial for sites
33 located at short distances from the rupture during large-magnitude earthquakes, where only
34 a relatively small number of records are currently available. The adoption of simulated
35 ground motions for engineering use, however, is only appropriate if they can adequately
36 reproduce the main characteristics of recorded ground motions. In the past decade, there
37 have been numerous efforts to validate simulated ground motions for engineering
38 applications (e.g., Bijelić et al., 2018; Burks et al., 2015; Burks and Baker, 2014; Fayaz et
39 al., 2020; Galasso et al., 2020; Teng and Baker, 2019). It is widely recognized that the
40 intensity of ground motions recorded during past earthquakes, typically characterized by
41 5%-damped response spectral ordinates, can greatly vary with changes in horizontal
42 orientation at a given site, in what is referred to as directionality (Hong and Goda, 2007;
43 Poulos and Miranda, 2022; Shahi and Baker, 2014). In this context, however, only a few
44 validation studies addressed directionality, and those that have done so have focused only
45 on the ratio between the maximum and median intensity (e.g., Burks et al., 2015; Burks
46 and Baker, 2014; Galasso et al., 2020; Teng and Baker, 2019). None have explicitly focused
47 on other directionality characteristics, such as the orientation of maximum horizontal
48 spectral response or its spatial distribution.

49 The lack of validation studies for the orientation of maximum response in simulated ground
50 motions may partially be attributed to the common belief that the maximum intensity
51 occurs at random, unpredictable orientations for source-to-site distances larger than 5 km.
52 The orientation of maximum response in recorded ground motions has historically been
53 studied by considering the angular distance of this orientation from the strike-normal
54 orientation. In their landmark paper, Somerville et al. (1997) observed that, for sites close
55 to faults and oscillators with periods longer than 0.6 s, the spectral accelerations in the
56 fault-normal orientation tend to be larger than average, which they attributed to radiation
57 patterns of shear waves. Since then, it is well documented that there is a higher probability
58 that the orientation of maximum response is closer to the strike-normal orientation for
59 rupture distances less than 5 km; however, at distances greater than 5 km, most studies

60 found that the maximum spectral response does not have a preferred orientation (Huang et
61 al., 2008; NEHRP Consultants Joint Venture, 2011; Shahi and Baker, 2014). Recently,
62 Poulos and Miranda (2023a) studied the orientation of maximum spectral response of
63 ground motions in the NGA-West2 database (Ancheta et al., 2014) by measuring and
64 conducting statistical analyses of the angular distance of this orientation from the epicentral
65 transverse orientation (i.e., the orientation of the ground motion transverse component,
66 which is perpendicular to a line connecting each recording station to the epicenter). Using
67 more than 5,000 recorded ground motions, they found that for ground motions recorded in
68 strike-slip earthquakes, there is a tendency for the orientation of maximum response to be
69 close to the epicentral transverse orientation and that, although less pronounced at shorter
70 periods, it gets closer to the transverse orientation as the oscillator period increases. These
71 new findings have also been independently verified in ground motions recorded during
72 recent strike-slip earthquakes not included in the NGA-West2 database (Girmay et al.,
73 2023, 2024; Poulos and Miranda, 2024).

74 Whilst there have been several studies on directionality of recorded earthquake ground
75 motions, there have been few on directionality of ground motion simulations, and none
76 specifically focusing on the orientation of maximum response. Burks and Baker (2014)
77 suggested a list of parameters that can be used as proxies to validate ground motion
78 simulations. In particular, to evaluate directionality, they recommended computing the ratio
79 of the maximum intensity (RotD100) and the median intensity of all orientations (RotD50)
80 for the simulations and comparing these ratios with those computed from recorded ground
81 motions and empirical models. Using a few records, they found that the variation of median
82 RotD100/RotD50 with changes in oscillator period for broadband simulations, although
83 generally matching the trend of recorded ground motions, tended to result in a more
84 polarized response for long-period oscillators. Similarly, Burks et al. (2015) found that the
85 RotD100 spectra of broadband simulations tended to be notably larger than empirical
86 ground motion prediction models (GMPMs) and that simulations tended to produce more
87 polarized motions compared to empirical predictions at long periods. They also found that
88 broadband simulations tended to produce ground motions with larger amplitudes in the
89 fault-normal orientations and lower amplitudes in the fault-parallel orientations when
90 compared to recorded motions, meaning that simulated ground motions are more polarized
91 than recorded ones. Other recent studies that have evaluated the possible use of simulated
92 ground motions for engineering applications have also found that the simulated motions,

93 on average, tended to produce larger RotD100/RotD50 ratios at long periods of vibration
94 when compared to recorded motions, and thus tended to polarize response more than
95 recordings (Galasso et al., 2020; Teng and Baker, 2019).

96 All studies identified above primarily focused on validating the level of polarization of the
97 simulated ground motions by using the RotD100/RotD50 ratio as a proxy and did not study
98 other directionality characteristics. This suggests a gap in existing literature regarding
99 directionality in simulated ground motions and, more specifically, the variation of the level
100 of intensity with changes in orientation or the orientation of maximum response. Therefore,
101 it is of interest to investigate these aspects of directionality in ground motion simulations,
102 particularly in the context of the recent findings by Poulos and Miranda (2023a) for strike-
103 slip earthquakes. At the center of Poulos and Miranda’s reasoning for investigating the
104 orientation of maximum response with respect to the epicentral transverse orientation is
105 that S waves from theoretical double couple point sources in a homogenous medium exhibit
106 polarization transverse to the direction of propagation (Aki and Richards, 2009). Thus,
107 evaluating whether this phenomenon is also apparent in physics-based simulations would
108 be important.

109 This work uses CyberShake Study 15.12 to evaluate directionality in simulated ground
110 motions from strike-slip earthquakes. CyberShake is a physics-based computational study
111 developed by the Southern California Earthquake Center (SCEC) to calculate the seismic
112 hazard for sites in California (Graves et al., 2011; Graves and Pitarka, 2010). This study
113 investigates whether the physics-based simulations sufficiently capture some directionality
114 characteristics observed in recorded earthquake ground motions. More specifically, the
115 orientation of maximum spectral response and the spatial distribution of response for 5%-
116 damped linear elastic oscillators when subjected to these simulated ground motions is
117 studied. The dependence of the orientation of maximum response on rupture variation and
118 oscillator period is investigated. The extent of polarization in the simulated ground motions
119 within the horizontal plane is studied for each period by evaluating spectral accelerations
120 in all non-redundant orientations.

121 **SELECTED GROUND MOTION SIMULATIONS**

122 The simulated ground motions used in this study were obtained from SCEC’s CyberShake
123 Study 15.12, which includes ground motion simulations for over 300 stations around the

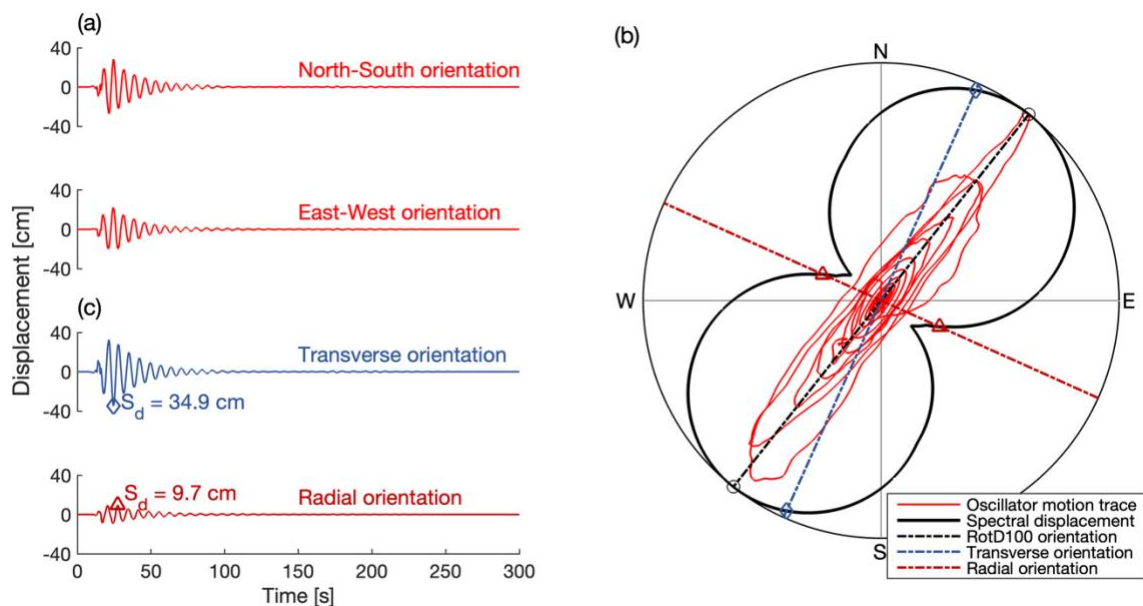
124 greater Los Angeles (LA) metropolitan region generated by considering all Uniform
125 California Earthquake Rupture Forecast, Version 2.0 (UCERF2.0) ruptures within 200 km
126 from each station (Graves et al., 2011). For each station, there are over 7,000 ruptures,
127 which in this context refers to a realization of an earthquake of a given magnitude, and
128 415,000 rupture variations, which refers to a realization of a particular hypocenter and slip
129 distribution within each rupture.

130 Specifically, this paper uses ground motions simulated at 334 stations from five different
131 rupture variations across two strike-slip ruptures, totaling 1,670 records. The two rupture
132 simulations studied include M_w 6.95 and M_w 7.45 events on the Elsinore fault zone. This
133 major active fault zone is part of the right-lateral strike-slip San Andreas fault system in
134 southern California, stretching over 200 km from the south-east boundary of the LA basin
135 to the border with Mexico (Hull and Nicholson, 1992; Magistrale and Rockwell, 1996).
136 Each of the two rupture simulations studied involve multiple segments of the fault zone,
137 with the smaller magnitude event rupturing the Whittier and Glen Ivy segments (W+GI)
138 and the larger magnitude event rupturing the Glen Ivy, Temecula, Julian, and Coyote
139 Mountain segments (GI+T+J+CM). To investigate the effect of hypocenter location and
140 direction of rupture propagation on the orientation of maximum intensity, different rupture
141 variations (i.e., hypocenter realizations) were chosen such that they initiate at distinctly
142 different hypocenter locations for each rupture. Specifically, for the W + GI rupture, these
143 include hypocenters at the south-most end (i.e., south-to-north rupture propagation), north-
144 most end (i.e., north-to-south rupture propagation), and center (i.e., bilateral rupture
145 propagation). For the GI+T+J+CM rupture, only hypocenters at the south-most and north-
146 most ends were considered. For each rupture variation described above, the velocity
147 seismograms at each station were downloaded and then numerically differentiated to obtain
148 acceleration waveforms, which were then used to compute the relative displacement
149 response of 5%-damped linear elastic single-degree-of-freedom oscillators. See the Data
150 and Resources section of this paper for the specific CyberShake database identifiers of the
151 ground motion simulations used in this study.

152 **ORIENTATION OF MAXIMUM SPECTRAL RESPONSE**

153 CyberShake simulated ground motions provide complete waveforms at a given station in
154 one vertical and two horizontal orientations, much like those in recorded ground motions.
155 In earthquake engineering, the most common measure of ground motion intensity is the

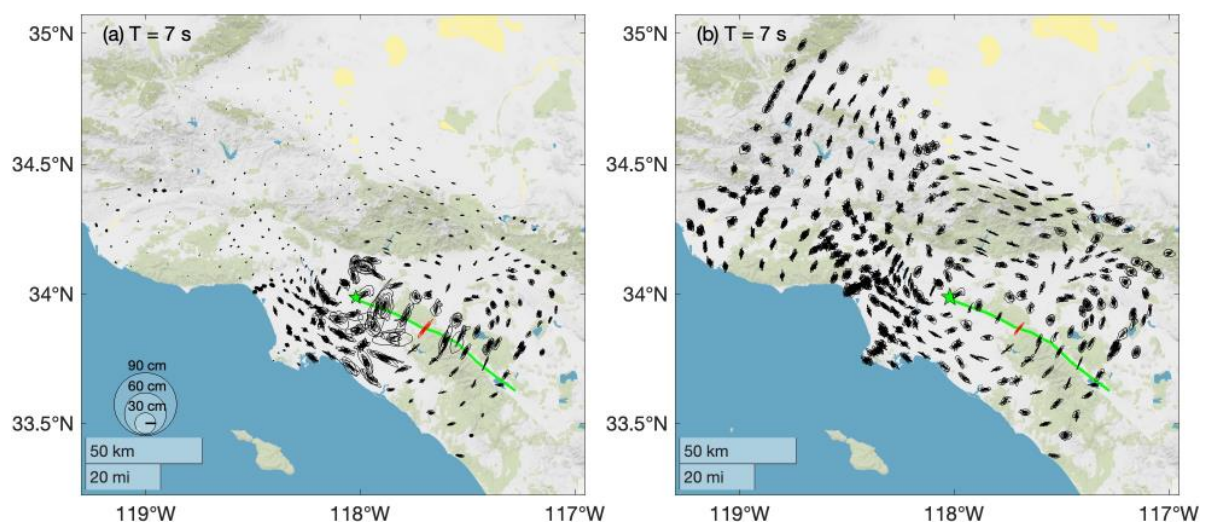
156 5%-damped response spectral ordinate, which represents the peak response of a single-
 157 degree-of-freedom linear-elastic oscillator with a 5% damping ratio. The availability of two
 158 orthogonal horizontal components at a station allows the computation of the bidirectional
 159 response of an oscillator within the horizontal plane, which represents the trace (or
 160 hodograph) of the oscillator's motion when subjected to the ground motion. For example,
 161 Figure 1a shows the relative displacement response of a 5%-damped linear-elastic 7 s
 162 oscillator subjected to the ground motion generated at station S597 from a simulation of a
 163 M_w 6.95 rupture on Elsinore fault (W + GI). Cross-plotting the response in the two
 164 orientations shown in Figure 1a results in the hodograph indicated by the red trace in Figure
 165 1b. Having the complete response in two orientations at a given station also allows for the
 166 computation of waveforms or spectral ordinates in any desired orientation. For instance,
 167 Figure 1c shows the relative displacement waveforms in the transverse and radial
 168 orientations.



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 170 **Figure 1.** An example of relative displacement response of a 7 s oscillator when subjected to the
 171 ground motion generated at station S597 from a simulation of a M_w 6.95 rupture on Elsinore fault
 172 (W + GI): (a) relative displacement response of the oscillator in the north-south and east-west
 173 orientations; (b) bidirectional response of the oscillator in the horizontal plane (i.e., hodograph) and
 174 the corresponding spectral displacement in all orientations, also indicating the peak relative
 175 displacement of the oscillator in the transverse and radial orientations; and (c) relative displacement
 176 response of the oscillator in the transverse and radial orientations. The circular marker indicates the
 177 maximum spectral response, and the diamond and triangular markers indicate the spectral ordinates
 178 in the transverse and radial orientations, respectively.

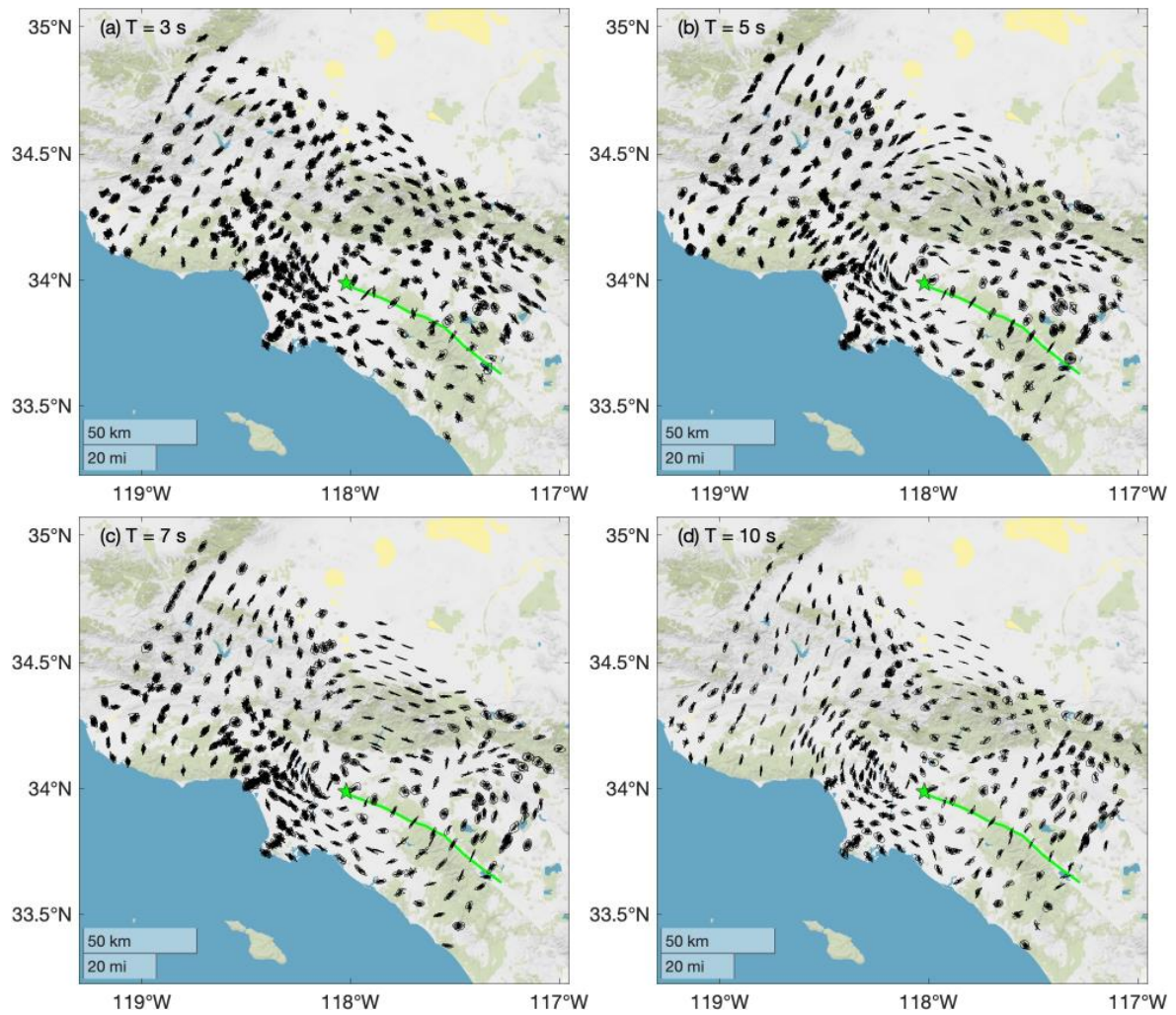
179 To get a better understanding of directionality in physics-based ground motion simulations,
 180 hodographs like that shown in Figure 1 can be computed and studied for all stations. Figure
 181 2a shows the spatial distribution of relative displacement hodographs of 5%-damped linear-

182 elastic 7 s oscillators subjected to all ground motions generated from a simulation of a M_w
183 6.95 rupture on Elsinore fault (W + GI). The hodograph in red corresponds to that for
184 station S597 (which was provided as an example in Figure 1), and the green line indicates
185 the rupture trace. The epicenter for the rupture is indicated by the green star. CyberShake
186 Study 15.12 was created to compute hazard curves at multiple sites in southern California,
187 which required generating waveforms at numerous stations with close spacing between
188 them. As a result, the spatial distribution of most stations shown in Figure 2a has a semi-
189 regular grid arrangement, making it particularly valuable to study directionality (in contrast
190 to real-world recording stations, which are often clustered around large urban areas and
191 their corresponding infrastructure). In Figure 2a, the size of each hodograph indicates the
192 station-to-station variation of intensity. It is apparent that the hodographs are largest at
193 stations close to the rupture and epicenter, and their size decreases with increasing rupture
194 distance. The large variation of intensity between stations makes it difficult to identify and
195 compare the shapes of the hodograph. In order to better visualize and compare both the
196 level of polarization and the orientation of maximum response at each station, the relative
197 displacement hodographs can be normalized by the maximum spectral response and scaled
198 to show the bidirectional motion at all stations with the same peak amplitude. For example,
199 Figure 2b shows the normalized version of Figure 2a, which permits a clearer comparison
200 of hodograph shapes for the event.



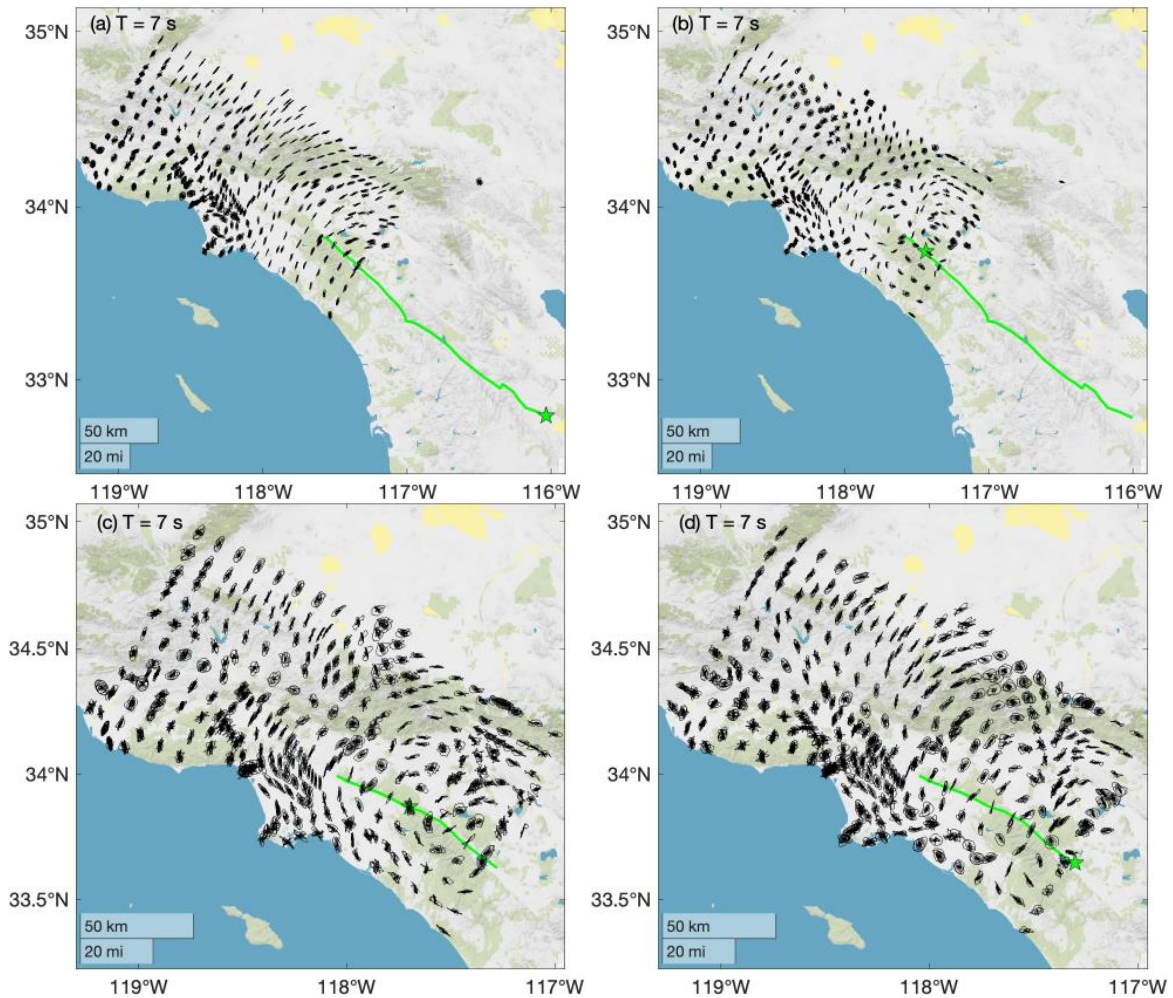
201 **Figure 2.** Spatial distribution of relative displacement hodographs of 5%-damped linear elastic
202 oscillators with periods $T = 7$ s subjected to simulated ground motions generated from a M_w 6.95
203 rupture on Elsinore fault (W + GI). The hodograph sizes are (a) relative, such that the size of each
204 hodograph indicates the station-to-station variation in the ground motion intensity, and (b)
205 normalized and scaled to show all stations with the same peak amplitude. The hodograph in red is
206 for station S597 (see Figure 1b).

207 Figure 3 shows the normalized relative displacement hodographs of 5%-damped linear-
208 elastic oscillators subjected to ground motions generated from the same rupture as that
209 shown in Figure 2, but for four different oscillator periods (3, 5, 7, and 10 s). In Figure 3,
210 stations near each other tend to have similar bidirectional responses, generally
211 characterized by being circumscribed by an approximate elliptical shape, regardless of the
212 oscillator period considered. As the oscillator period increases, its response becomes
213 “cleaner” and more elliptical (i.e., with notably larger major axes relative to minor axes) at
214 all distances from the rupture. This implies that there is an increase in the level of
215 polarization in the oscillator’s bidirectional response as its period of vibration increases.
216 Additionally, adjacent stations tend to have similar orientations of the principal axes of the
217 ellipse, suggesting that the orientation of maximum response is not entirely random and
218 that the motions have some preferred orientation of maximum response. The bidirectional
219 responses in Figure 3 are consistent with those observed in oscillators subjected to recorded
220 earthquake ground motions from strike-slip earthquakes, as noted by Poulos and Miranda
221 (2023a) and others (Girmay et al., 2023, 2024).



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Figure 3. Normalized relative-displacement hodographs of 5%-damped linear elastic oscillators subjected to ground motions generated from a simulation of a M_w 6.95 rupture on Elsinore fault (W + GI) for oscillators having periods of (a) $T = 3$ s, (b) $T = 5$ s, (c) $T = 7$ s, and (d) $T = 10$ s. The epicenter is indicated by the green star, and the rupture propagates from north to south along the fault trace shown by the green line.



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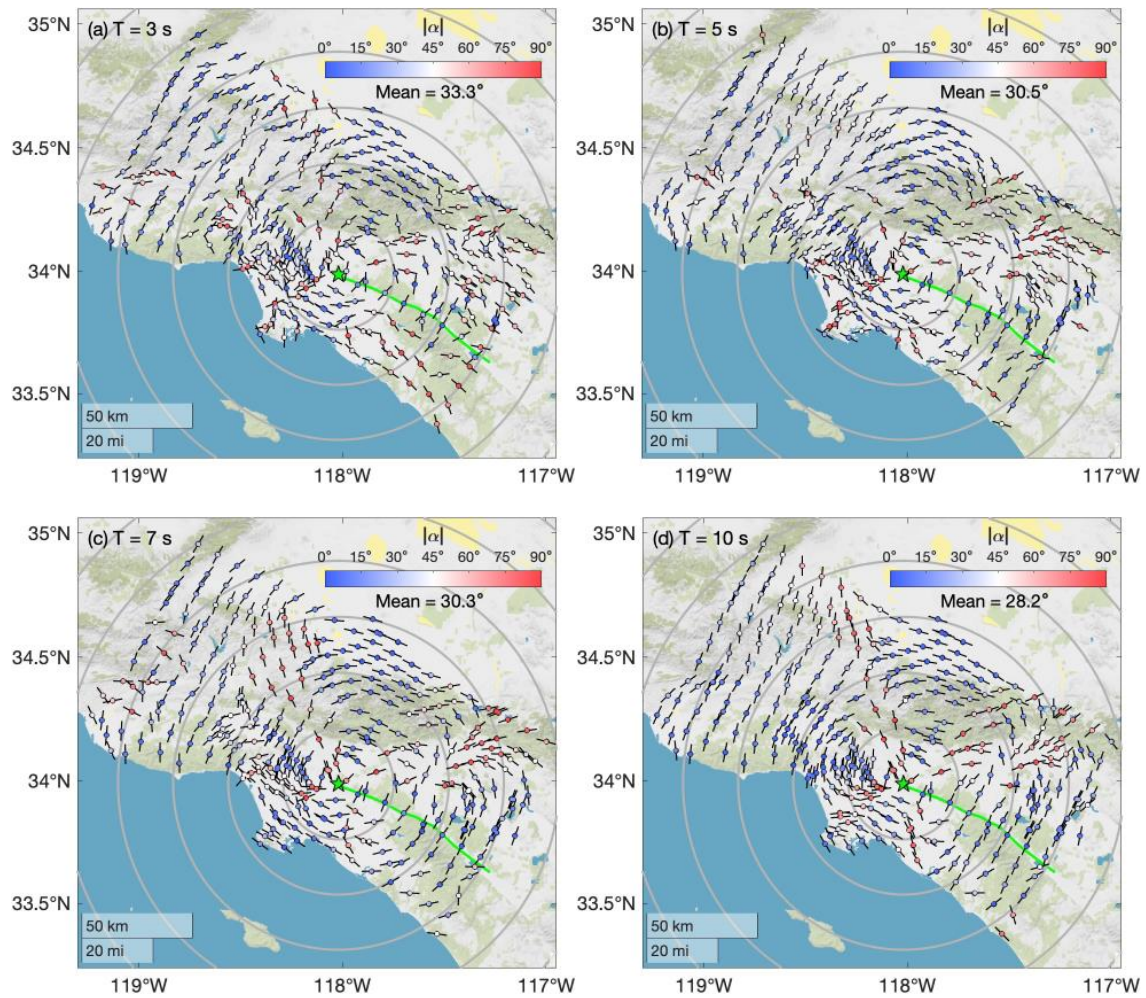
Figure 4. Normalized relative-displacement hodographs of 5%-damped linear elastic oscillators with periods $T = 7$ s subjected to ground motions generated from simulation of ruptures on Elsinore fault with (a) $M_w = 7.45$ and rupture running south-north, (b) $M_w = 7.45$ and rupture running north-south, (c) $M_w = 6.95$ and bilateral rupture, and (d) $M_w = 6.95$ and rupture running south-north. Ruptures in (a) and (b) consist of GI + J + T + CM segments, while ruptures in (c) and (d) consist of W + GI segments. The green line indicates the rupture trace, and the epicenter is indicated by the green star.

Should these observations result from a physical phenomenon, it would be expected that similar observations would hold for both recorded and simulated ground motions, and additionally for other realizations of simulated ruptures with the same faulting mechanism. Figure 4 shows the bidirectional response of 7 s oscillators subjected to simulated ground motions generated due to four different rupture scenarios on the Elsinore fault. The surface projection of the rupture trace for each scenario is again indicated by the green line, and the green star indicates the epicenter. The bidirectional oscillator responses to ground motions from each rupture simulation show trends similar to those observed in Figure 3. Most stations tend to have an elliptical response, and the response of adjacent stations remains similar regardless of the earthquake's magnitude, rupture length, or location of the

246 epicenter. The stability of this observation across different strike-slip rupture realizations
247 is consistent with that observed for ground motions recorded during strike-slip earthquakes.
248 For example, Poulos and Miranda (2023a) observed similar bidirectional responses for
249 strike-slip earthquakes in the NGA-West2 database, such as the 1991 M_w 7.1 Hector Mine
250 earthquake. More recently, Girmay et al. (2023, 2024) noted similar observations for the
251 2022 M_w 6.4 Ferndale earthquake and the 2023 Türkiye earthquake doublet, respectively.
252 For a record that has waveforms available in two orthogonal horizontal orientations, the
253 two waveforms can be combined to compute the peak oscillator response at any specific
254 azimuth and rotated incrementally to obtain the variation of spectral ordinates in each
255 nonredundant orientation (i.e., a range of azimuths of 180° , see Figure 1b) (Boore, 2010).
256 Within the nonredundant orientations, the 100th percentile of spectral response is referred
257 to as RotD100 (also known as the maximum spectral response), and the 50th percentile is
258 referred to as RotD50 (also known as the median spectral response) (Boore, 2010). Thus,
259 the orientation of maximum spectral response corresponds to the azimuth at which the
260 RotD100 intensity occurs while the RotD50 orientation can occur at two or more different
261 orientations. Visually, the maximum spectral response at each station shown in Figures 2 –
262 4 corresponds to the point in each hodograph that is furthest away from the resting position
263 of the oscillator (as illustrated in Figure 1b).

264 As discussed in the introduction, most, if not all, prior studies that have studied the
265 directionality of simulated ground motions have focused on whether the ratio between the
266 maximum response and the median response with the horizontal plane is consistent with
267 that computed from recorded ground motions. The median of this ratio has a stable and
268 known variation with changes in the period of vibration of the oscillator (Burks and Baker,
269 2014), and thus is often used as a proxy in ground motion simulations validation exercises
270 to evaluate directionality (Burks et al., 2015; Burks and Baker, 2014; Galasso et al., 2020;
271 Teng and Baker, 2019). More recent findings by Poulos and Miranda (2023a) have
272 identified that for strike-slip earthquakes, the orientation of maximum spectral response
273 tends to be close to the transverse orientation. They found that this observation is less
274 pronounced at short periods but becomes apparent with increasing period. Therefore, it
275 would be important to verify whether the simulated ground motions exhibit orientations of
276 maximum response that are consistent with those observed in recorded ground motions.

277 In their study of the orientation of maximum spectral response in recorded ground motions,
278 Poulos and Miranda (2023a) defined the angle between the orientation of RotD100 and the
279 epicentral transverse orientation (i.e., the orientation orthogonal to a line connecting the
280 epicenter to the station) as $\alpha \in [-90^\circ, 90^\circ]$. Measured relative to the transverse
281 orientation, this angle is negative if RotD100 orientation is clockwise or positive if
282 counterclockwise. A similar measure of angular difference between the transverse and
283 RotD100 is $|\alpha| \in [0^\circ, 90^\circ]$, which does not distinguish between the clockwise or
284 counterclockwise orientation of RotD100 with respect to the transverse orientation. In this
285 study, distinguishing between clockwise or counterclockwise is not considered since,
286 according to Poulos and Miranda (2023a), the mean variation of intensity is symmetric with
287 respect to the orientation of RotD100. Therefore, the angular difference $|\alpha|$ is used.



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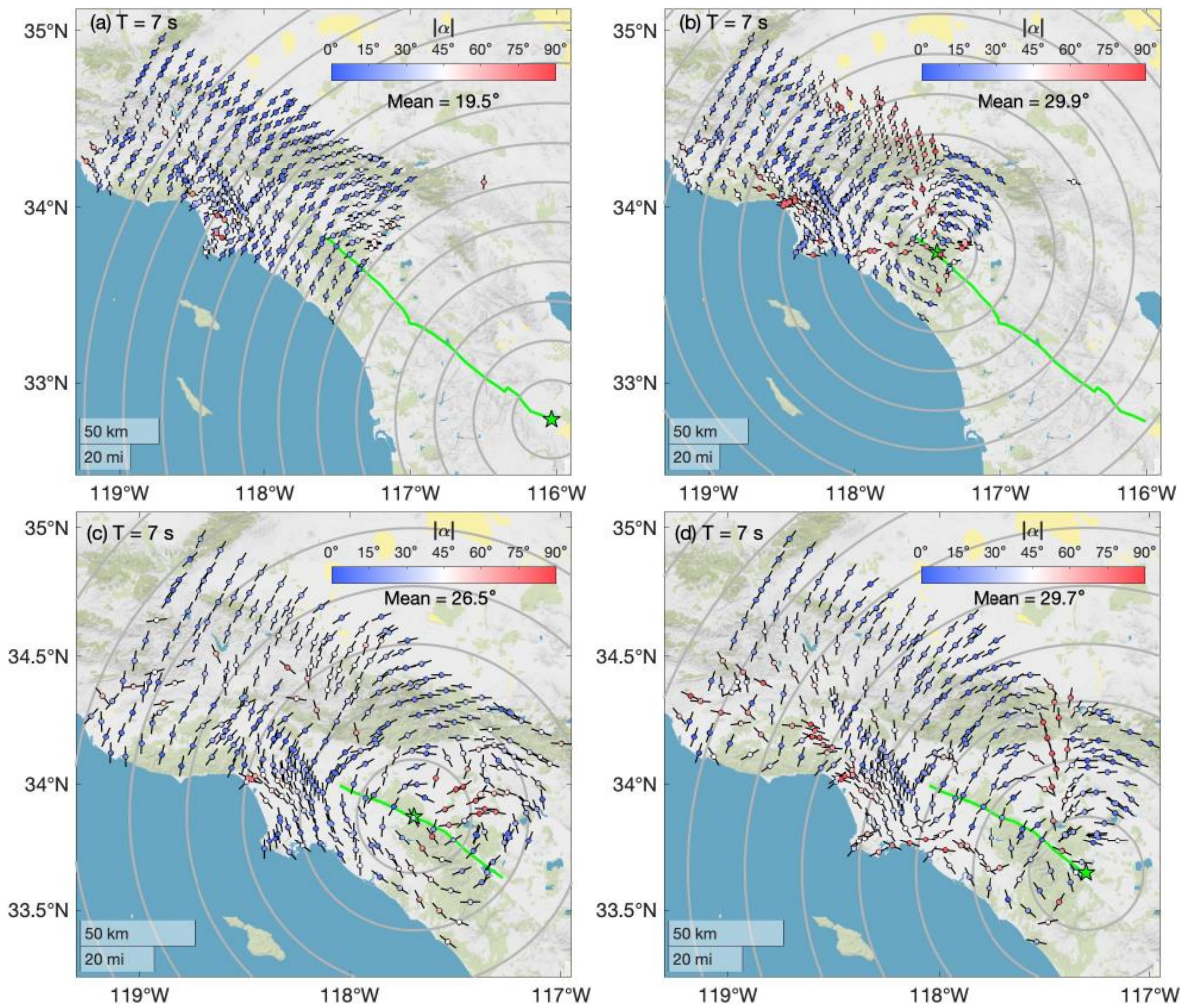
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Figure 5. Orientation of maximum spectral response of 5%-damped linear elastic oscillators with periods of (a) 3 s, (b) 5 s, (c) 7 s, and (d) 10 s, as indicated by the short black line segments at each station (indicated by a circle), and their angular difference from the transverse orientation, as indicated by the color inside each circle. Each station is subjected to ground motions generated from a simulation of a M_w 6.95 rupture on the Elsinore fault (W + GI). The large concentric grey circles centered around the epicenter indicate epicentral transverse orientations.



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Figure 6. Effect of rupture variation on the orientation of maximum spectral response of 5%-damped linear elastic oscillators with periods of 7 s. The orientation of maximum spectral response at each station is indicated by the short black line segments at each station, and the angular difference from the transverse orientation is indicated by the color inside each circle. The large concentric grey circles centered around the epicenter indicate the epicentral transverse orientations.

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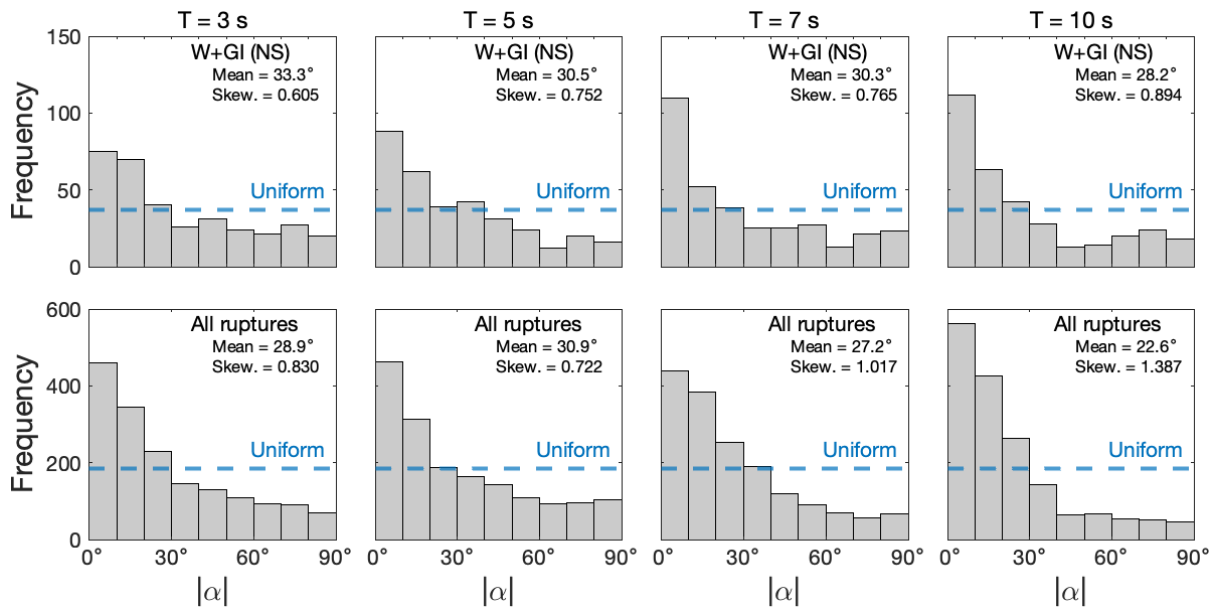
Figure 5 uses short black line segments to indicate the RotD100 orientation at each station for oscillators with four different fundamental periods subjected to ground motions generated from a simulated M_w 6.5 earthquake on the Elsinore fault (W + GI). The color inside the circles at each station indicates the angular difference between the RotD100 orientation and the epicentral transverse orientation, with red tones signifying RotD100 orientations closer to the radial orientation and blue tones signifying RotD100 orientations closer to the transverse orientation. Generally, the orientations of the short black line segments in Figure 5 are typically similar for adjacent stations and, for most stations, tend to form a circular pattern around the epicenter. As the period of vibration of the oscillators increases, the circular pattern around the epicenter becomes more apparent. Most stations

311 also have circles with blue tones, suggesting that, for most stations, the RotD100
312 orientation occurs close to the transverse orientation across all four periods shown in the
313 figure. Furthermore, as the oscillator period increases, more stations become blue colored,
314 indicating that more stations have RotD100 orientations that occurred close to the
315 transverse orientation. To better understand the effects of rupture variations, Figure 6 shows
316 the spatial distribution of $|\alpha|$ but for four different rupture realizations on the Elsinore fault
317 for a single oscillator period of 7 s. Much like what is observed in Figure 5, most stations
318 in Figure 6 also have blue-toned circles, suggesting that the RotD100 orientation remains
319 close to the transverse orientation regardless of rupture length, rupture initiation, or rupture
320 direction. These observations are consistent with those made from recorded earthquake
321 ground motions.

322 Prior studies on ground motions recorded from recent strike-slip earthquakes (Girmay et
323 al., 2023, 2024; Poulos and Miranda, 2023a) have found that the empirical probability
324 distribution of $|\alpha|$ deviates significantly from a uniform distribution and has a predictable
325 pattern, with larger density occurring at lower $|\alpha|$ values. Therefore, should the simulated
326 ground motions appropriately capture directionality characteristics, then the empirical
327 distribution of $|\alpha|$ for the simulated ground motions should follow the same, or at least a
328 very similar, pattern as that from recorded ground motions. While Figures 5 and 6 use
329 colors to indicate the spatial distribution of $|\alpha|$ in the simulated ground motions, Figure 7
330 presents the empirical probability distribution of $|\alpha|$ for four oscillators subjected to ground
331 motions generated from different earthquake ruptures simulated on the Elsinore Fault. The
332 top row shows probability distributions of $|\alpha|$ for four different periods of vibration
333 computed from a single realization that ruptures the Elsinore fault from north to south
334 involving the Whittier and Glen Ivy segments; similarly, the second row aggregates the
335 probability distribution of $|\alpha|$ considering all five rupture variations considered in this
336 study for a total 1,670 simulated ground motions. Inset in each panel of the figure is the
337 mean $|\alpha|$ and total number of records used for each case.

338 Figure 7 shows that the empirical distribution of $|\alpha|$ is skewed towards small values of $|\alpha|$
339 for all oscillator periods, with the greatest density between 0° and 30° . As the oscillator
340 period increases, the distribution gets more skewed, as indicated by the increased density
341 at lower $|\alpha|$ and by the increase in the skewness coefficient. This observation is true for
342 both the single M_w 6.95 rupture on the W+GI segments and for the aggregated data using

343 simulated ground motions from all ruptures. If RotD100 was equally likely in any
 344 orientation with respect to the transverse orientation, then the observed distribution would
 345 have been uniform as indicated by the dashed blue horizontal line, and the mean $|\alpha|$
 346 be 45° . However, the distributions are highly skewed and the mean $|\alpha|$ for each period is
 347 notably below 45° , indicating that the orientation of RotD100 tends to occur close to the
 348 epicentral transverse orientation. These empirical probability distributions are consistent
 349 with the findings for ground motions recorded during strike-slip earthquakes, suggesting
 350 that the physics-based simulations generally capture the main features of the directionality
 351 effects present in recorded ground motions.



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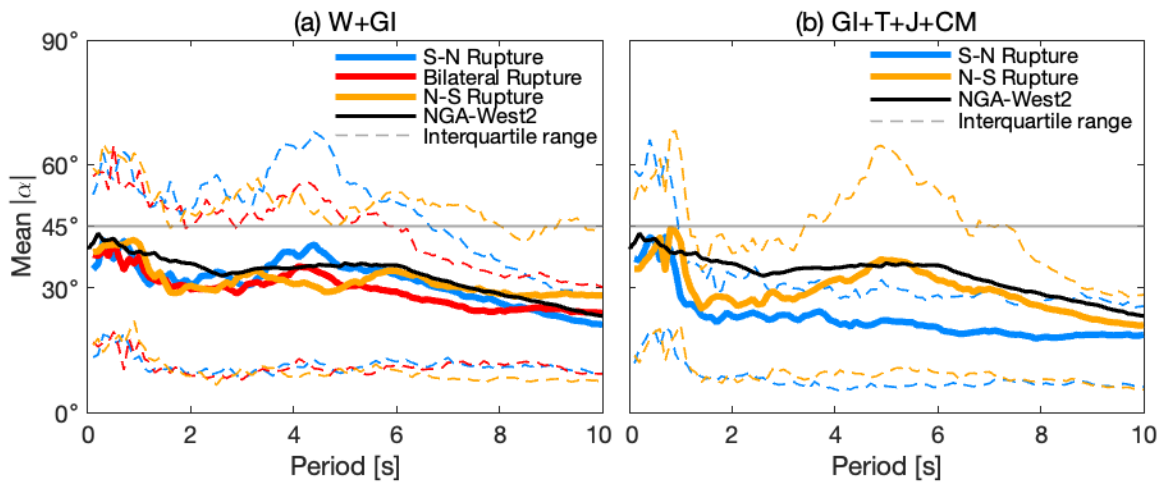
353 **Figure 7.** Histograms of the angular difference between the transverse orientation and the RotD100
 354 orientation for oscillators subjected to simulated ground motions from different rupture variations
 355 simulated on the Elsinore Fault. The top row presents results for a simulated M_w 6.95 event on the
 356 W + GI segments rupturing north-to-south (NS), whilst the bottom row presents results computed
 357 from simulations of five different realizations, comprised of three M_w 6.95 ruptures on the W + GI
 358 segments and two M_w 7.45 ruptures on the GI + J + T + CM segments. The dotted line represents
 359 the histogram should the orientation of RotD100 be equally likely to occur at any orientation with
 360 respect to the transverse orientation. Each panel indicates the mean $|\alpha|$ and the skewness
 361 coefficient. A total of 334 records were used for the top row, and 1,670 were used for the bottom
 362 row.

363 As previously mentioned, one trend that has been found in ground motions recorded during
 364 strike-slip earthquakes is the tendency for the mean values of $|\alpha|$ to decrease with
 365 increasing oscillator period. In other words, in recorded strike-slip earthquakes, it has been
 366 found that the orientation of RotD100 tends to get closer to the transverse orientation as
 367 the period of vibration of the oscillator increases. This is often attributed to wave scattering

368 since higher frequencies have shorter wavelengths, which makes them more susceptible to
369 heterogeneities in the Earth's crust. To check whether this trend is observed and is similar
370 in the physics-based ground motion simulations, Figure 8 plots the variation of mean $|\alpha|$
371 as a function of oscillator period for five different rupture variations on the Elsinore fault.
372 Figure 8a presents the results for variations that ruptured the W + GI segments, whilst 8b
373 shows the results for variations that ruptured the GI+T+J+CM segments. Also included in
374 the figure and indicated by the black line is the mean $|\alpha|$ for 1,962 ground motions recorded
375 during strike-slip earthquakes from the NGA-West2 database as computed by Poulos and
376 Miranda (2023a). There are three main observations to note from this figure. First, the
377 variation of mean $|\alpha|$ on different variations of the smaller rupture (W + GI) is more similar
378 to each other than it is in the larger rupture (GI+T+J+CM). Second, mean $|\alpha|$ follows
379 similar trends for different rupture realizations, where the RotD100 orientation generally
380 gets closer to the transverse orientation as the period of vibration of the oscillator increases,
381 suggesting that the descending trend is a stable characteristic regardless of the direction of
382 rupture propagation. Third, the mean $|\alpha|$ is consistently below 45° for all rupture variations,
383 suggesting that the orientation of RotD100 tends to occur relatively close to the transverse
384 orientation, which is consistent with the behavior observed in recorded ground motions in
385 the NGA-West2 database. This is further highlighted in Figure 9, which indicates that the
386 probability that $|\alpha|$ falls between 0° and 30° is more than double than that of falling
387 between 30° and 60° or between 60° and 90° for most oscillator periods. The findings in
388 Figures 8 and 9 generally follow the trends observed for real recorded ground motions by
389 Poulos and Miranda (2023a) and Girmay et al. (2023), suggesting that the physics-based
390 simulations exhibit directionality effects that, in general, are consistent with those observed
391 in recorded ground motions.

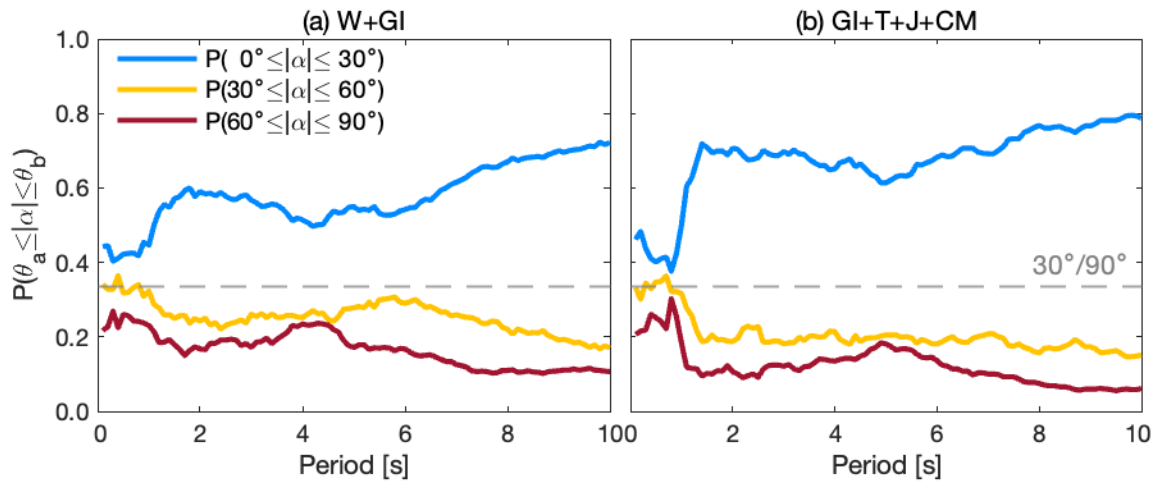
392 It is important to note that the significant decrease in the probability that the RotD100
393 intensity occurs between 0° and 30° from the transverse orientation for periods of vibration
394 between 0 – 1 s, followed by a rapid increase after 1 s, is more accentuated in the simulated
395 ground motions when compared to that observed in recorded ground motions. This may be
396 attributed to the hybrid broadband method used to generate the waveforms, which
397 combines deterministic and semi-stochastic approaches. More specifically, the method for
398 simulation (Graves et al., 2011; Graves and Pitarka, 2010) is deterministic for frequencies
399 below 1 Hz and incorporates a theoretically robust representation of fault rupture and wave

400 propagation. For frequencies above 1 Hz, the method uses a stochastic representation of
 401 source radiation combined with a simplified theoretical representation of wave propagation
 402 and scattering effects (Graves and Pitarka, 2010), which seems to adequately capture the
 403 probabilities in Figure 9 given the similarity to the probabilities reported by Girmay et al.
 404 (2023). Above 2 s, the probabilities are consistent with real ground motions, where the
 405 physics dominates (i.e., the polarization of S-waves that result in the orientation of
 406 RotD100 being close to the transverse orientation is apparent, and wave scattering effects
 407 are not as significant). However, the rapid increase in the probability that RotD100 in
 408 simulated ground motions falls between 0° and 30° from the transverse orientation between
 409 1 s and 2 s may be artificially large due to the transition between the stochastic and
 410 deterministic approaches.



411

412 **Figure 8.** Effect of the period of vibration of the oscillator on the mean angular difference between
 413 the orientation of RotD100 and the transverse orientations for five different rupture variations on
 414 the Elsinore fault's (a) W + GI segments, and (b) GI + T + J + CM segments. The solid lines present
 415 the mean angular difference whilst the dashed lines present the interquartile range. Results for
 416 variations that ruptured south-to-north (S-N) are shown in blue, while yellow is used for variations
 417 that ruptured north-to-south (N-S). The black line shows the mean angular difference for strike-slip
 418 records in the NGA-West2 database, as computed by Poulos and Miranda (2023a).



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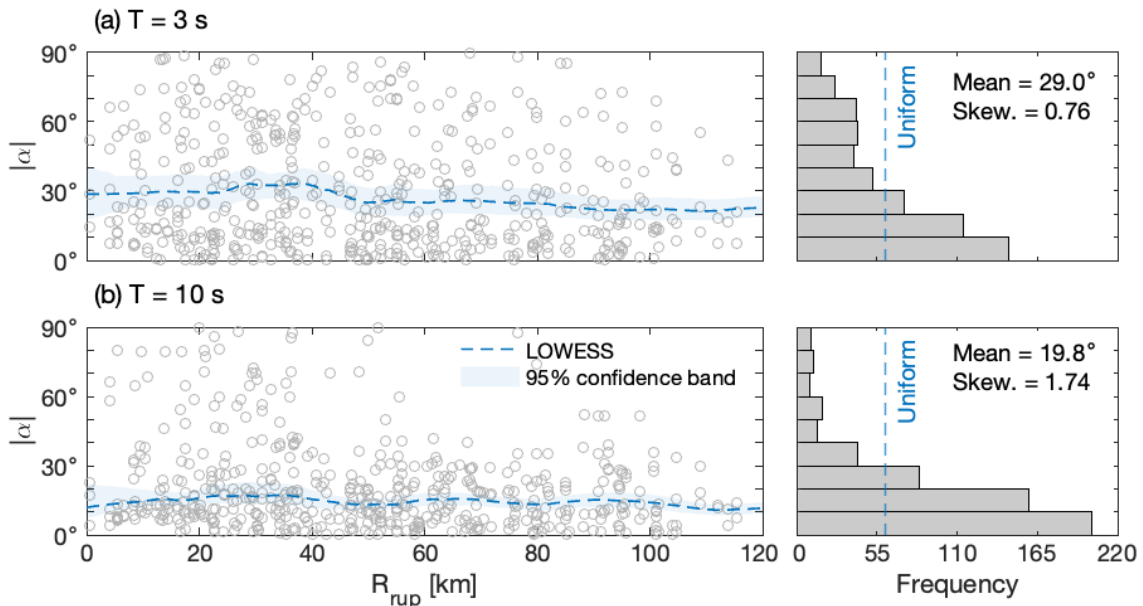
442

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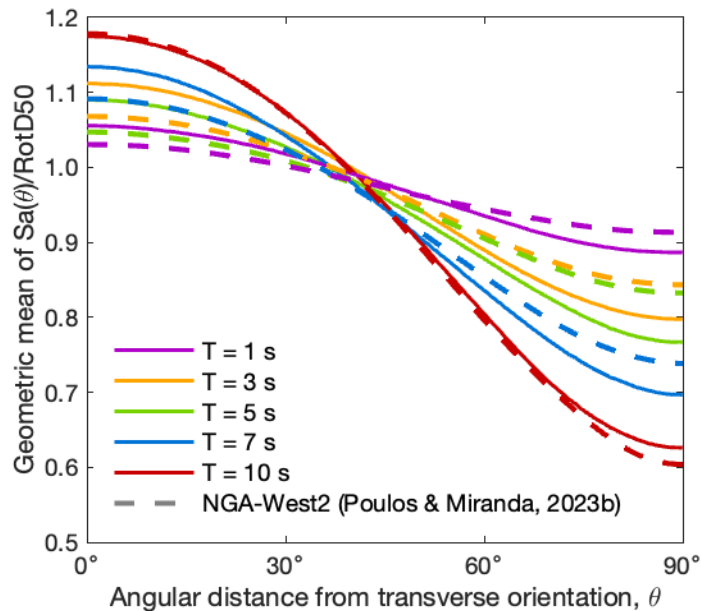
Figure 9. Probability of the angular difference between the orientation of RotD100 and the transverse orientation occurring between $[0^\circ, 30^\circ]$, $[30^\circ, 60^\circ]$, and $[60^\circ, 90^\circ]$ for all rupture variations considered in this study on Elsinore fault segments (a) W + GI, and (b) GI + T + J + CM. While studying ground motion directionality of spectral ordinates in the 2023 Türkiye earthquake, Girmay et al. (2023) found that the mean $|\alpha|$ at a given period remains relatively constant for Joyner-Boore distances up to the maximum studied (around 400 km). Similar findings were observed for the 2022 Ferndale earthquake, where the epicentral distance was found to have a minimal effect on the distribution of $|\alpha|$ with the exception of the near field region (Girmay et al., 2024). This suggests that an approximately constant mean $|\alpha|$ across a wide range of distances is a stable characteristic of ground motion directionality in recorded motions, so it is important to verify whether simulated ground motions also exhibit this trend.

Figure 10 shows the possible effect of rupture distance on $|\alpha|$ for oscillators with periods of 3 s and 10 s, located at stations where rupture distance information is available. The locally weighted scatter plot smoothing (LOWESS) curve, which is indicated by the blue line, represents the change in mean $|\alpha|$ with increasing rupture distance. On the right-hand side of the figure is the corresponding empirical distribution of $|\alpha|$ and its mean. This figure shows that the LOWESS curve for each oscillator period remains relatively constant across all rupture distances, suggesting that the rupture distance does not substantially influence the orientation of RotD100 with respect to the transverse orientation. Additionally, the scatter points tend to be more concentrated towards the bottom third of the figure, which is further indicated by the histograms on the right that get more skewed with increasing period, signifying that the orientation of RotD100 gets closer to the transverse orientation with increasing period. These observations and trends are consistent with those observed

444 in recorded ground motion, suggesting that the physics-based simulations exhibit realistic
 445 directionality characteristics.



446
 447 **Figure 10.** Evaluation of the possible influence of rupture distance on the angular difference
 448 between the orientation of RotD100 and the transverse orientation for (a) $T = 3$ s, and (b) $T = 10$ s.
 449 The dashed lines represent the locally weighted scatterplot smoothing (LOWESS), and the shaded
 450 band represents the 95% confidence band. The results presented consider all stations across all
 451 rupture variations considered in this study that had rupture distances available in the CyberShake
 452 study database (a total of 550 records). For each oscillator period, histograms of the angular
 453 difference between the transverse and RotD100 orientations are presented on the right.



454
 455 **Figure 11.** Effect of the angular distance from the transverse orientation on the geometric mean
 456 ratio of the spectral response at a given angle from the transverse orientation and the RotD50
 457 intensity for all rupture variations considered in this study. The solid lines represent geometric
 458 means of the simulated ground motions for different oscillator periods, whilst the dashed lines
 459 represent the results for ground motions recorded during strike-slip earthquakes in the NGA-West2
 460 database (Poulos and Miranda, 2023b).

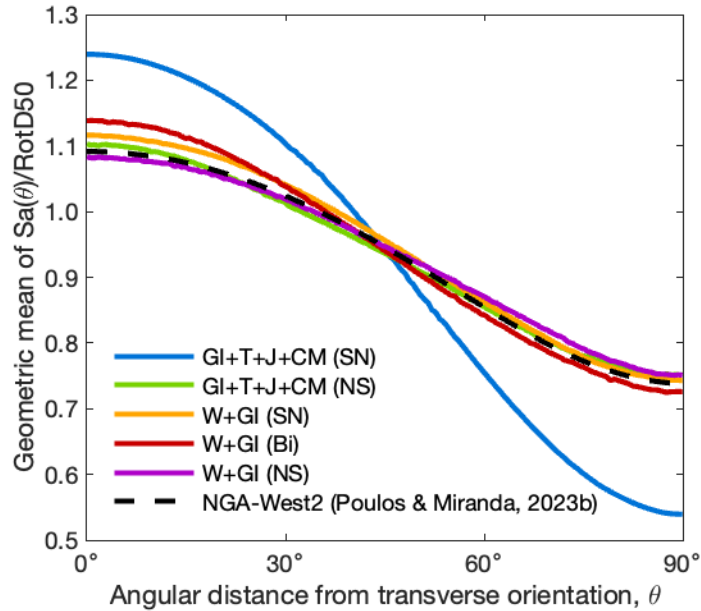
461 **ORIENTATION-SPECIFIC SPECTRAL ACCELERATIONS**

462 In the previous section, it has been shown that the orientation of RotD100 for physics-based
463 simulations tends to occur close to the epicentral transverse orientation, which is consistent
464 with observations from recorded ground motions. However, another aspect of directionality
465 that has not been previously studied for physics-based simulated ground motions is the
466 variation of ground motion intensity with changes in orientation within the horizontal
467 plane. So, it is important to study this and compare the findings to those from recorded
468 ground motions.

469 Poulos and Miranda (2023b) found that orientation-specific horizontal spectral
470 accelerations for ground motions recorded from strike-slip earthquakes have generally
471 stable characteristics when computed as a function of the angular distance from the
472 transverse orientation and normalized by the associated orientation-independent median
473 intensity (RotD50). Here, their approach is used to investigate whether the orientation-
474 specific spectral accelerations for simulated ground motions exhibit a similar variation to
475 that observed in recorded ground motions when computed as a function of the angular
476 distance from the transverse orientation.

477 Figure 11 shows the effect of angular distance from the transverse orientation on the
478 geometric mean ratio of spectral response at a given angle θ from the transverse and the
479 RotD50 intensity (i.e., $Sa(\theta)/RotD50$). The solid lines indicate the geometric mean of
480 $Sa(\theta)/RotD50$ for different oscillator periods subjected to simulated ground motions for
481 all rupture variations considered in this study. Also included in the figure and indicated by
482 the dashed lines are the corresponding geometric mean of $Sa(\theta)/RotD50$ for 1,962 ground
483 motions recorded during strike-slip earthquakes from the NGA-West2 database (Poulos
484 and Miranda, 2023b). Figure 11 shows that the general shape of $Sa(\theta)/RotD50$ for the
485 simulated ground motions follows the general trend of recorded ground motions. At long
486 periods, and more specifically at 10 s, the geometric mean of $Sa(\theta)/RotD50$ for the
487 simulated ground motions agrees closely with that of the NGA-West2 records, especially
488 at angles within 60° from the transverse orientation. With decreasing periods, the geometric
489 mean of $Sa(\theta)/RotD50$ somewhat deviates from that for the recorded ground motions,
490 although the general shape is maintained. The difference between the simulated and
491 recorded ground motions in the short period range may be attributed to wave scattering.
492 Specifically, there is more wave scattering at higher frequencies (i.e., short periods), which

493 may not be well captured in the physics-based simulations, which use a semi-stochastic
 494 method coupled with simplified scattering effects in the high frequencies. However, in the
 495 lower frequencies (i.e., as the oscillator period increases), wave scattering effects become
 496 less important, and physics-based simulations exhibit directionality characteristics
 497 consistent with those of recorded ground motions.



498

499 **Figure 12.** Geometric mean ratio of the spectral response at a given angle from the transverse
 500 orientation and the RotD50 intensity for 7 s oscillators as a function of the angular distance from
 501 the transverse orientation. The solid lines represent different rupture variations, while the dashed
 502 line shows the results for ground motions recorded during strike-slip earthquakes in the NGA-
 503 West2 database (Poulos and Miranda, 2023b).

504 A measure of the level of polarization at different periods can be determined from Figure
 505 11 by comparing the ratio of $Sa(\theta)/RotD50$ at $\theta = 0^\circ$ and $\theta = 90^\circ$. This ratio
 506 approximates how much larger the intensity in the major response axis is compared to that
 507 in the minor response axis. Therefore, the larger this ratio, the greater the level of
 508 polarization in the oscillator response. Figure 11 shows that, except for oscillators with
 509 periods of 10 s, the simulated ground motions have a geometric mean of $Sa(0^\circ)/RotD50$
 510 greater than those of records in the NGA-West2 database, whilst also having
 511 $Sa(90^\circ)/RotD50$ less than those of records in the NGA-West2. This indicates that the
 512 simulated ground motions have a level of polarization that is, on average, greater than that
 513 observed in recorded ground motions.

514 Figure 12 shows the geometric mean of $Sa(\theta)/RotD50$ of 7 s oscillators as a function of
 515 angular distance from the transverse orientation for each simulated rupture variation. Also

516 shown is the geometric mean for ground motions recorded during strike-slip earthquakes
517 in the NGA-West2 database. From this figure, it is apparent that, in general, the
518 $Sa(\theta)/RotD50$ ratio is relatively consistent regardless of rupture variation and generally
519 agrees well with real ground motions in the NGA-West2 database. However, one notable
520 exception is the GI+T+J+CM (SN) rupture realization, which exhibits a much larger level
521 of polarization compared to the NGA-West2. The higher level of polarization in this case
522 can be explained by the location of most stations relative to the location of the epicenter in
523 that particular rupture. More specifically, the epicenter for the GI+T+J+CM (SN) rupture
524 (see Figure 6a) occurs such that the azimuth of each station relative to the epicenter is
525 similar for all stations, which results in the transverse orientation at each station generally
526 coinciding. Ultimately, this results in most stations having very similar levels of response
527 polarization, which also produces a higher mean polarization. Nonetheless, the results in
528 Figures 11 and 12 show that the physics-based simulations, despite resulting in a somewhat
529 larger level of polarization than expected, still have variations in intensity with changes in
530 orientation that are consistent with those of recorded ground motions.

531 **SUMMARY AND CONCLUSIONS**

532 This study used simulated ground motions obtained from CyberShake Study 15.12 to
533 investigate several aspects of the directionality in simulated ground motions. The
534 observations for simulated ground motions were compared to those from ground motions
535 recorded during strike-slip earthquakes. Linear elastic oscillators at more than 300
536 locations were subjected to simulated ground motions from five rupture realizations. Then,
537 the following aspects were carefully studied: (1) variation of the level of directionality with
538 changes in the period of the oscillators; (2) possible effect of distance to the source on the
539 level of polarization; (3) orientation of maximum oscillator (i.e., spectral) response and its
540 spatial distribution; and (4) variation of the level of ground motion intensity as one rotates
541 away from the transverse orientation. Additionally, the possible effects of rupture variation
542 (i.e., rupture length, point of initiation of the rupture, and direction of rupture propagation)
543 on directionality were investigated and compared to trends observed in recorded ground
544 motions.

545 The orientation of maximum response of 5%-damped single-degree-of-freedom oscillators
546 subjected to ground motions from physics-based simulated strike-slip ruptures tends to
547 occur systematically close to the epicentral transverse orientation, consistent with recent

548 findings for records obtained from strike-slip earthquakes. Oscillator hodographs at
549 stations close to each other were found to be similar over the spatial region that was studied.
550 The hodographs at neighboring stations tended to become even more similar and polarized
551 with increasing period. In addition to the increase in polarization, the orientation of
552 maximum intensity got closer to the epicentral transverse orientation as the oscillator
553 period increased. The orientation of maximum spectral response remained close to the
554 transverse orientation for all rupture variations considered. This is consistent with previous
555 observations for ground motions recorded during strike-slip earthquakes with magnitudes
556 larger than or equal to five.

557 Lastly, orientation-specific spectral accelerations normalized by the median intensity from
558 all orientations of the simulated ground motions, when computed as a function of angular
559 distance from the transverse orientation, were generally found to exhibit variations
560 consistent with records in the NGA-West2 database. At long periods, where the wave
561 scattering effects are less important, the physics-based simulations exhibit mean variation
562 of intensity as a function of the angular distance from the transverse orientation that closely
563 aligns with that from recorded ground motions. However, at short periods, where the
564 physics-based simulations use semi-stochastic methods and simplified scattering effects,
565 the simulated waveforms exhibit a level of polarization that is greater than that observed in
566 recorded ground motions, suggesting that ground motions simulated with the hybrid
567 approach are not fully able to reproduce all the wave scattering that occurs near the surface.
568 This might be because the crustal structure used in the CyberShake Study 15.12 did not
569 consider shear wave velocities lower than 500 m/s.

570 Ground motion simulations have the potential to play an important role in earthquake
571 engineering and design, especially in the absence of suitable real records (such as for short
572 source-to-site distances and large magnitudes). While numerous prior validation studies
573 focused on various features of these simulated ground motions, none solely focused on
574 directionality. The findings in this paper demonstrate that the physics-based ground
575 motions for strike-slip earthquakes, despite their higher-than-expected level of polarization
576 at short periods, sufficiently capture stable directionality characteristics observed in
577 recorded ground motions. These results provide more confidence in using simulated ground
578 motions for appropriate engineering applications, particularly for structures with periods
579 longer than 1 s where directionality effects are more important.

580 **DATA AND RESOURCES**

581 All simulated ground motions used in this study were obtained from CyberShake Study
 582 15.12 (https://strike.scec.org/scecpedia/CyberShake_Study_15.12). Unique database
 583 identifiers for each rupture realization used are summarized below. Map figures in this
 584 paper used tiles by MapTiler and base-map data by OpenStreetMap.

Fault	Source_ID	Rupture_ID	Rup_Var_ID	ERF_ID	Rup_Var_Scenario_ID	SGT_Variation_ID
Elsinore (W+GI) S-N Rupture	17	0	0	36	6	8
Elsinore (W+GI) Bilateral	17	0	10			
Elsinore (W+GI) N-S Rupture	17	0	17			
Elsinore (GI+T+J+CM) S-N Rupture	10	0	0			
Elsinore (GI+T+J+CM) N-S Rupture	10	0	168			

585

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600 **DECLARATION OF CONFLICTING INTEREST**

601 The authors declare no conflicts of interest with respect to the research, authorship, and/or
602 publication of this article

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